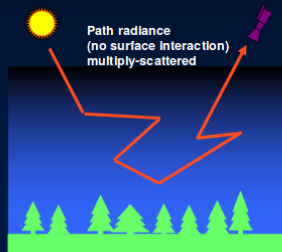
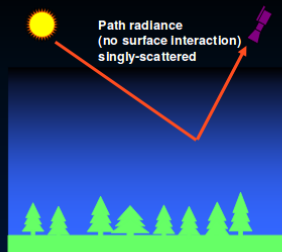
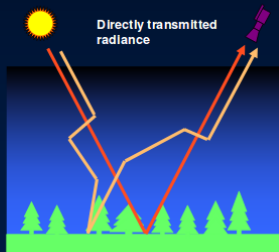


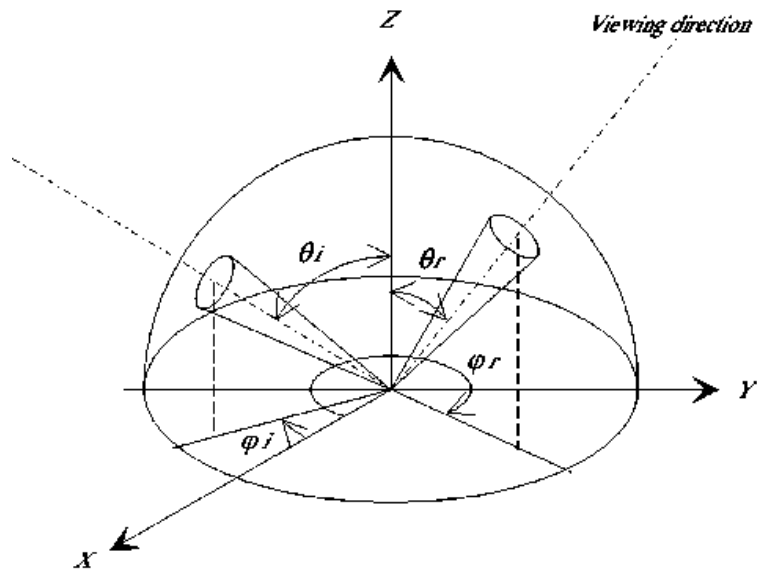
Sensori per aerosol a bordo di satelliti

- ▶ Lo studio di aerosol da satellite ha avuto inizio dalla necessità di correggerne l'effetto sulle misure di Ozono e della riflettività superficiale
- ▶ Nel caso dell'ozono l'assorbimento è largo, per cui l'estinzione da aerosol può contaminare la misura
- ▶ Nel visibile-NIR la radianza dovuta a riflessioni dalla superficie è dello stesso ordine di grandezza (tranne casi particolari) della radianza dovuta all'atmosfera



Top-of-atmosphere radiation consists of singly- and multiply-scattered components that may or may not have interacted with the surface





Formula per la radianza misurata

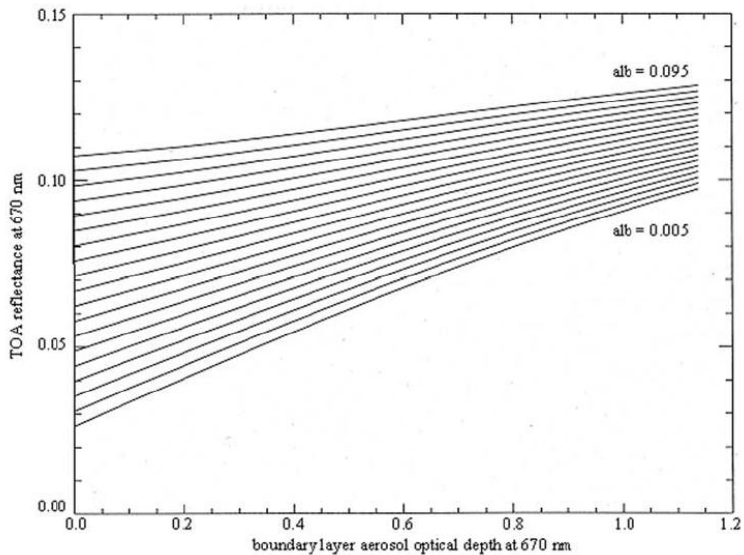
$$L_0(\mu, \phi) = L_0(0, \mu, \phi) + \frac{RF_0(\tau, \mu)}{1 - RS}$$

- ▶ L_0 path radiance
- ▶ $F_0(\tau, \mu)$ flusso solare alla superficie
- ▶ R riflettanza della superficie
- ▶ S flusso riflesso dall'atmosfera

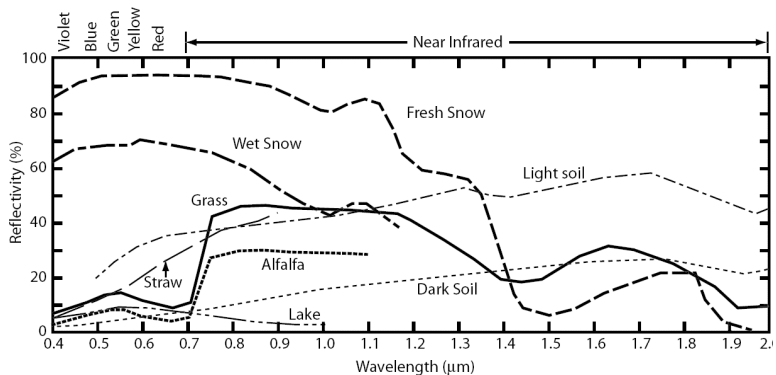
In prima approssimazione la path radiance é data dal singolo scattering, per spessori ottici piccoli:

$$L_0 = \frac{\omega_0}{4\pi} F_0 P(\Theta) \frac{\tau^*}{\mu}$$

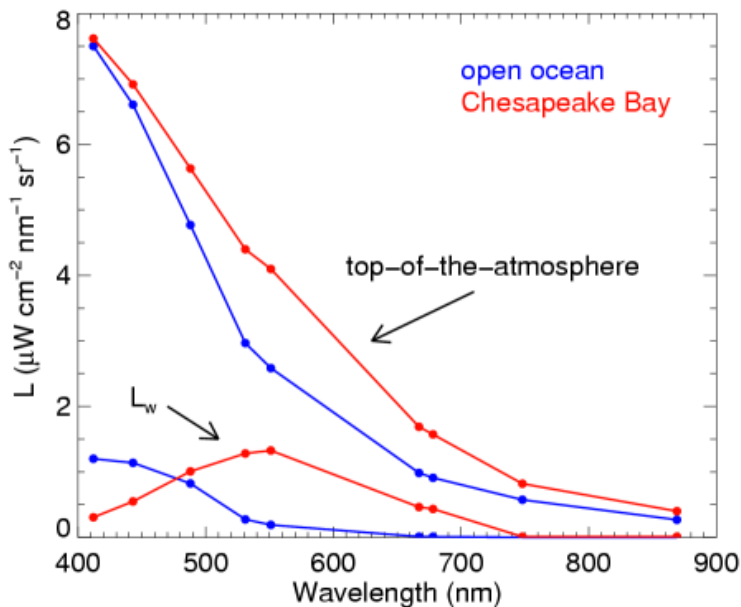
Quindi L_0 dipende sia dall'angolo di scattering che dallo spessore ottico



Riflettanze superficiali



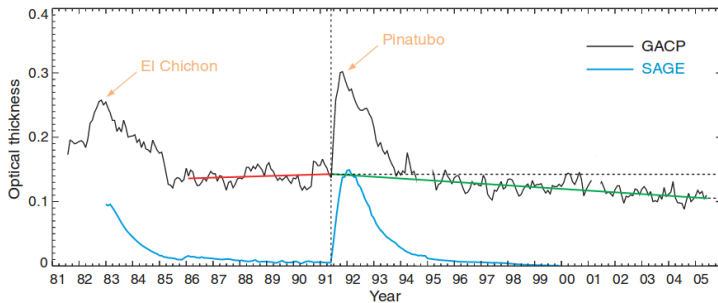
Riflettanze dal mare



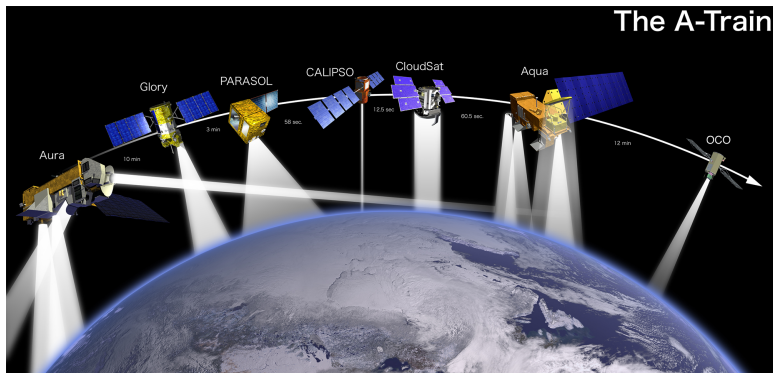
Misure di spessore ottico

- ▶ É necessario conoscere la riflettanza superficiale
- ▶ É necessario ipotizzare il tipo di aerosol
- ▶ Si può ottenere una misura affidabile solo quando la riflettanza superficiale é piccola (*dark target*, in pratica sul mare nel vicino IR)

Serie a lungo termine di spessori ottici



Satelliti dedicati allo studio dell'atmosfera: A-train



Su ogni satellite sono imbarcati sensori multispettrali attivi e passivi per le diverse zone spettrali.

I dati sono usati per una caratterizzazione completa dell'atmosfera e della superficie.

Sensori dedicati anche agli aerosol

MODIS : MODerate resolution Imaging Spectroradiometer

Sul satellite Aqua e Terra

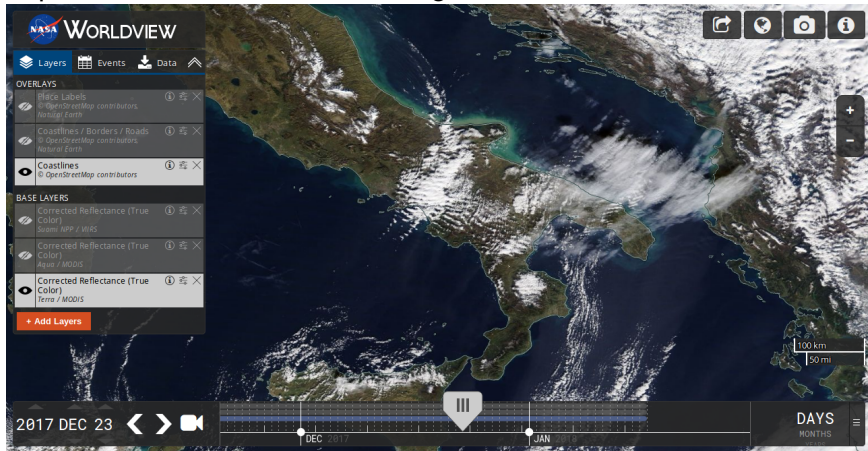
36 bande spettrali da visibile a IR

MISR : Multi-angle Imaging Spectro-Radiometer Spectrometer
(Su Terra)

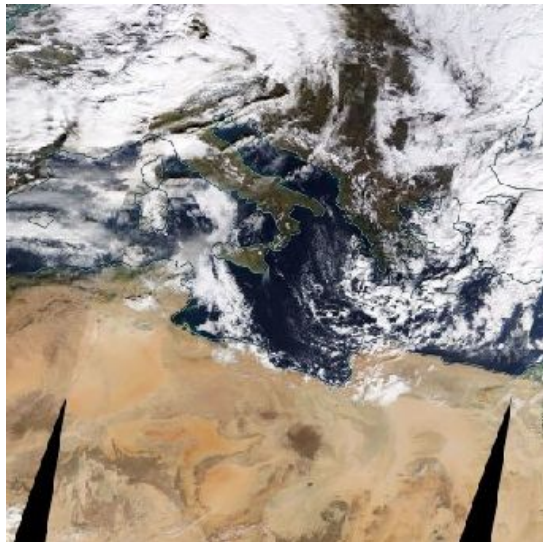
4 lunghezze d'onda e 9 angoli di scattering.

Immagini quasi in tempo reale

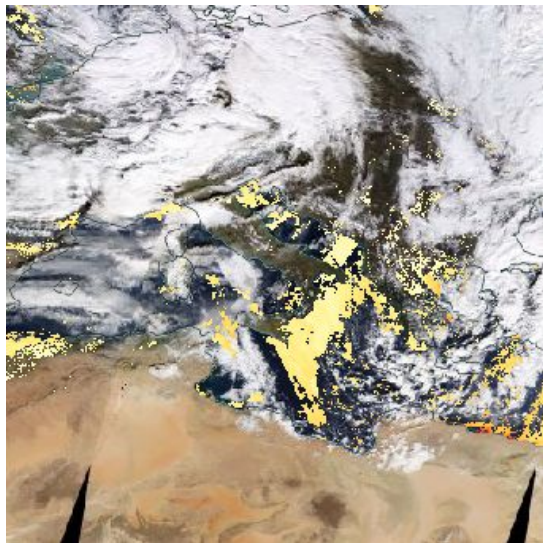
<https://worldview.earthdata.nasa.gov>



Immagini quasi in tempo reale



Immagini quasi in tempo reale



MODIS, Dust over Sahara,

(R: $0.66\mu\text{m}$, G: $0.55\mu\text{m}$, B: $0.47\mu\text{m}$)



Jan. 7, 2002 (007.1125)

(R: $2.13\mu\text{m}$, G: $1.64\mu\text{m}$, B: $1.24\mu\text{m}$)



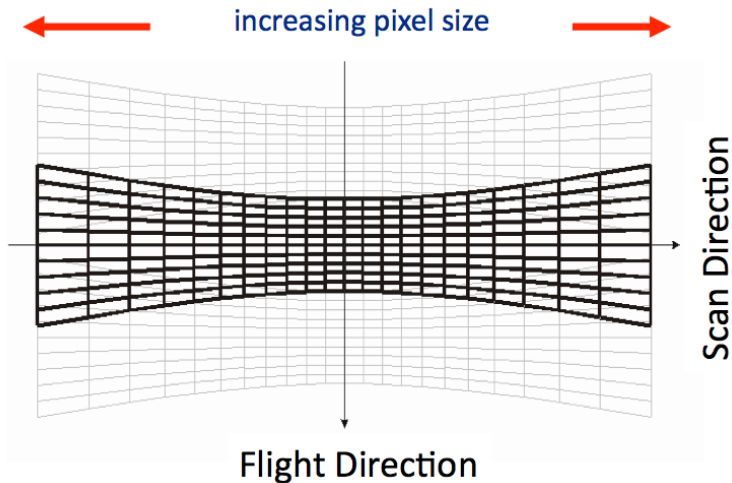
MODIS, Smoke over Australia, Dec. 25, 2001 (359.2345)

(R: $0.66\mu\text{m}$, G: $0.55\mu\text{m}$, B: $0.47\mu\text{m}$)

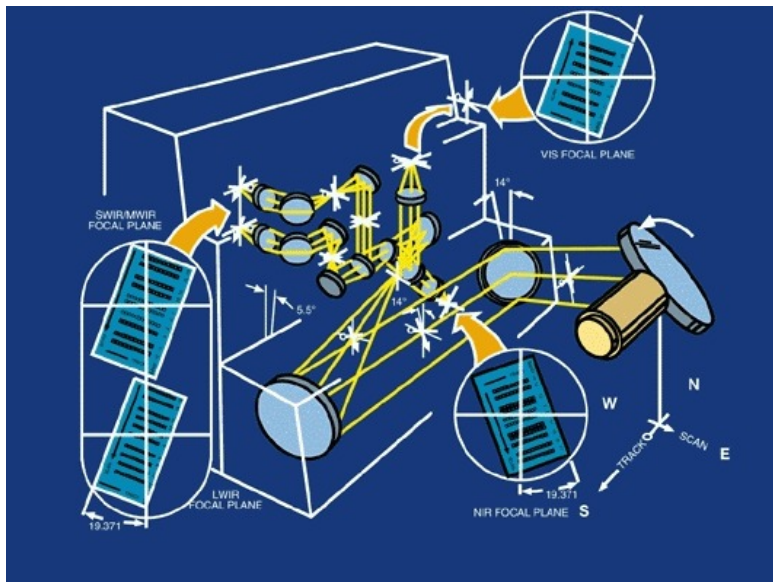


(R: $2.13\mu\text{m}$, G: $1.64\mu\text{m}$, B: $1.24\mu\text{m}$)



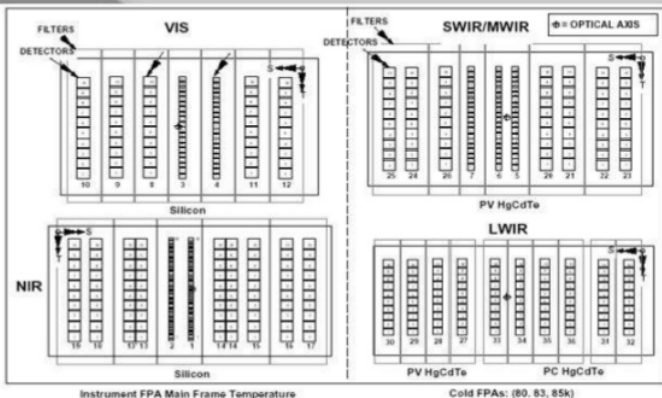


pixel size: 250 m - 1 km (dipende dalla lunghezza d'onda)
larghezza scansione (swath): 2300 km





MODIS FOCAL PLANE



Bande spettrali di MODIS - solari

Primary Use	Band No.	Bandwidth (nm)	Spectral Radiance	Required SNR
Land/Cloud Boundaries	1**	620-670	21.8	128
	2**	841-876	24.7	201
Land/Cloud Properties	3*	459-479	35.3	243
	4*	545-565	29.0	228
	5*	1230-1250	5.4	74
	6*	1628-1652	7.3	275
	7*	2105-2155	1.0	110
Ocean Color/ Phytoplankton/ Biogeochemistry	8	405-420	44.9	880
	9	438-448	41.9	838
	10	483-493	32.1	802
	11	526-536	27.9	754
	12	546-556	21.0	750
	13	662-672	9.5	910
	14	673-683	8.7	1087
	15	743-753	10.2	586
	16	862-877	6.2	516
Atmospheric Water Vapor	17	890-920	10.0	167
	18	931-941	3.6	57
	19	915-965	15.0	250

Bande spettrali di MODIS - IR termico

Primary Use	Band	Bandwidth (μm)	Spectral Radiance	Required NEDT (K)
Surface/Cloud Temperature	20	3.660-3.840	0.45(300K)	0.05
	21	3.929-3.989	2.38(335K)	2.00
	22	3.929-3.989	0.67(300K)	0.07
	23	4.020-4.080	0.79(300K)	0.07
Atmospheric Temperature	24	4.433-4.498	0.17(250K)	0.25
	25	4.482-4.549	0.59(275K)	0.25
Cirrus Clouds Water Vapor	26	1.360-1.390	6.00	150 (SNR)
	27	6.535-6.895	1.16(240K)	0.25
	28	7.175-7.475	2.18(250K)	0.25
	29	8.400-8.700	9.58(300K)	0.05
Ozone	30	9.580-9.880	3.69(250K)	0.25
Surface/Cloud Temperature	31	10.780-11.280	9.55(300K)	0.05
	32	11.770-12.270	8.94(300K)	0.05
Cloud Top Altitude	33	13.185-13.485	4.52(260K)	0.25
	34	13.485-13.785	3.76(250K)	0.25
	35	13.785-14.085	3.11(240K)	0.25
	36	14.085-14.385	2.08(220K)	0.35

Spectral Radiance values are in $\text{W/m}^2\text{-}\mu\text{m-sr}$

NEDT = Noise-equivalent temperature difference

Come si ottengono le proprietà degli aerosol?

- ▶ Occorre fare un'ipotesi sulla riflettività del terreno (si distinguono i casi di riflettività dal mare e dal terreno. Si escludono i casi con nuvole e superfici brillanti)
- ▶ Si riduce la complessità degli aerosol a un certo numero di tipi (size distribution e indice di rifrazione)
- ▶ Si variano le combinazioni tra i vari tipi di aerosol in modo da riprodurre al meglio le radianze osservate.
- ▶ risultato: spessore ottico a 0.55 nm e frazione del modo fine.
- ▶ Sul terreno si ottiene anche la riflettività a 2.5μ

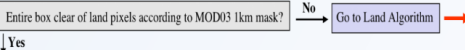
Proprietà degli aerosol usate per MODIS

F	$\lambda=0.47\text{--}0.86\mu\text{m}$	$\lambda=1.24\mu\text{m}$	$\lambda=1.64\mu\text{m}$	$\lambda=2.12\mu\text{m}$	r_g	σ	r_{eff}	Comments
1	1.45-0.0035i	1.45-0.0035i	1.43-0.01i	1.40-0.005i	0.07	0.40	0.10	Water Soluble
2	1.45-0.0035i	1.45-0.0035i	1.43-0.01i	1.40-0.005i	0.06	0.60	0.15	Water Soluble
3	1.40-0.0020i	1.40-0.0020i	1.39-0.005i	1.36-0.003i	0.08	0.60	0.20	Water Soluble with humidity
4	1.40-0.0020i	1.40-0.0020i	1.39-0.005i	1.36-0.003i	0.10	0.60	0.25	Water Soluble with humidity

C	$\lambda=0.47\text{--}0.86\mu\text{m}$	$\lambda=1.24\mu\text{m}$	$\lambda=1.64\mu\text{m}$	$\lambda=2.12\mu\text{m}$	r_g	σ	r_{eff}	Comments
5	1.35-0.001i	1.35-0.001i	1.35-0.001i	1.35-0.001i	0.40	0.60	0.98	Wet sea salt type
6	1.35-0.001i	1.35-0.001i	1.35-0.001i	1.35-0.001i	0.60	0.60	1.48	Wet sea salt type
7	1.35-0.001i	1.35-0.001i	1.35-0.001i	1.35-0.001i	0.80	0.60	1.98	Wet sea salt type
8	1.53-0.003i (0.47) 1.53-0.001i (0.55) 1.53-0.000i (0.66) 1.53-0.000i (0.86)	1.46-0.000i	1.46-0.001i	1.46-0.000i	0.60	0.60	1.48	Dust-like type
9	1.53-0.003i (0.47) 1.53-0.001i (0.55) 1.53-0.000i (0.66) 1.53-0.000i (0.86)	1.46-0.000i	1.46-0.001i	1.46-0.000i	0.50	0.80	2.50	Dust-like type

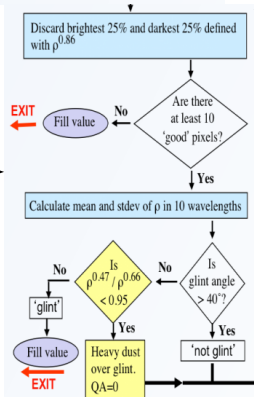
MODIS Aerosol Over Ocean Algorithm – 10 km product

All procedures applied to individual boxes of 20×20 pixels at 500 m resolution (10 km at nadir)



Continue masking pixel by pixel

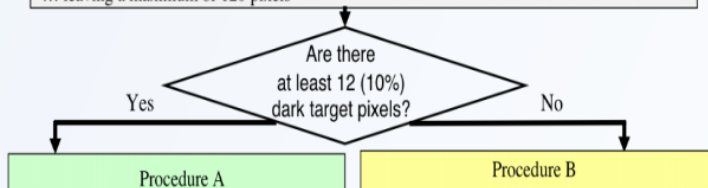
- (1) spatial variability: stdev of $\rho^{0.55}$ calculated from 3×3 centered on pixel > 0.0025 then all 9 pixels labeled 'cloudy'.
- (2) dust call back: if $\rho^{0.47}/\rho^{0.66} < 0.75$, then dust, use even if variability is high.
- (3) Threshold test: if $\rho^{0.47} > 0.40$ then 'cloudy'
- (4) IR tests: if any of 3 specific MOD35 tests indicate 'cloudy', then 'cloudy'
- (5) 1.38 cirrus tests: cloudy if $\rho^{1.38} > 0.03$
 OR if $0.005 \leq \rho^{1.38}/\rho^{1.24} \leq 0.30$
 AND $\rho^{1.38} > 0.03$
 but, if $0.10 \leq \rho^{1.38}/\rho^{1.24} \leq 0.30$
 AND $0.01 \leq \rho^{1.38} \leq 0.03$
 'not cloudy' but QC=0
- (6) sediment mask: Described by Li et al., 2003

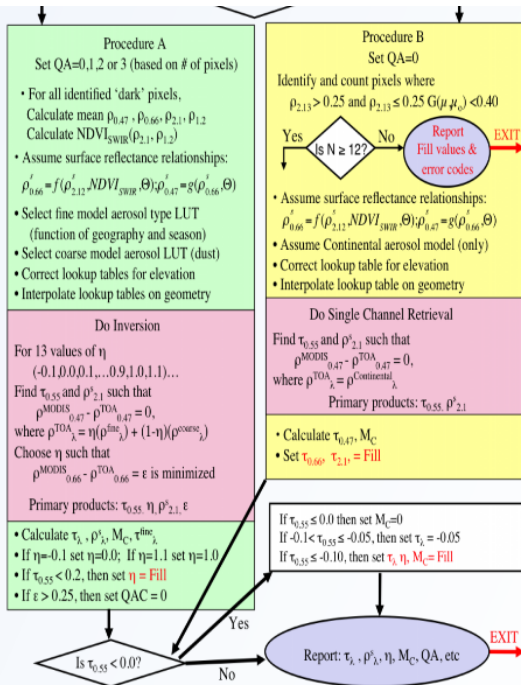


MODIS Over Land Algorithm

All procedures applied to individual boxes of 20 x 20 pixels at 500 m resolution (10 km at nadir)

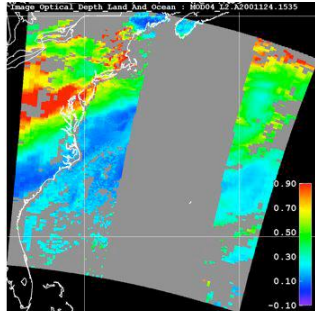
- Ensure angles and reflectance values are valid. If not: **report Fill values and EXIT**
- Identify and mask (discard) all water, cloudy and snow/ice pixels.
- Identify “dark target pixels” that have $0.01 \leq \rho_{2.13} \leq 0.25$
- Discard brightest 50% and darkest 20% of pixels defined with $\rho_{0.66}$
... leaving a maximum of 120 pixels



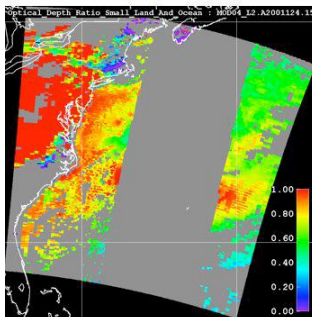




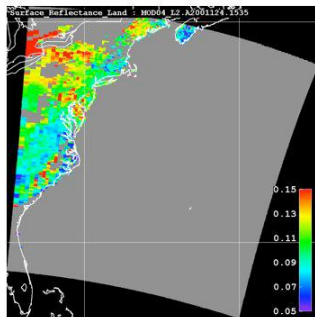
A) RGB: $\rho^M(0.66, 0.55, 0.47 \mu\text{m})$



B) $\tau(0.55 \mu\text{m})$



C) $\eta(0.55 \mu\text{m})$



D) $\rho^s(2.1 \mu\text{m})$

9 view angles at Earth surface with 14-bit pushbroom cameras

7 minutes to view each scene from all 9 angles

275 m spatial resolution per pixel

~400-km swath width

Calibrated measurements of the intensity of reflected sunlight

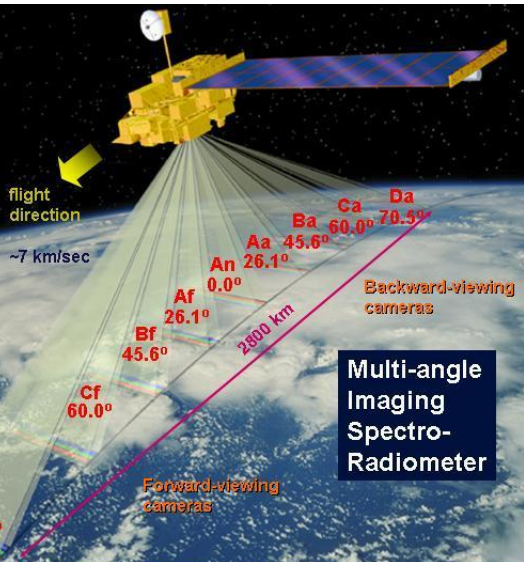
4 spectral bands at each angle:

446 nm \pm 21 nm

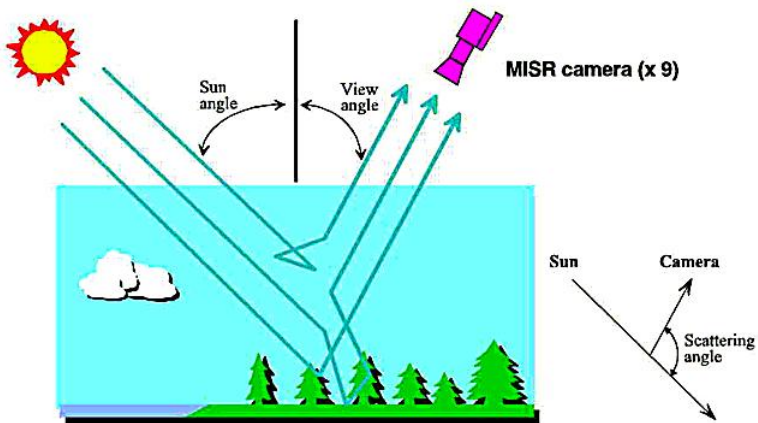
558 nm \pm 15 nm

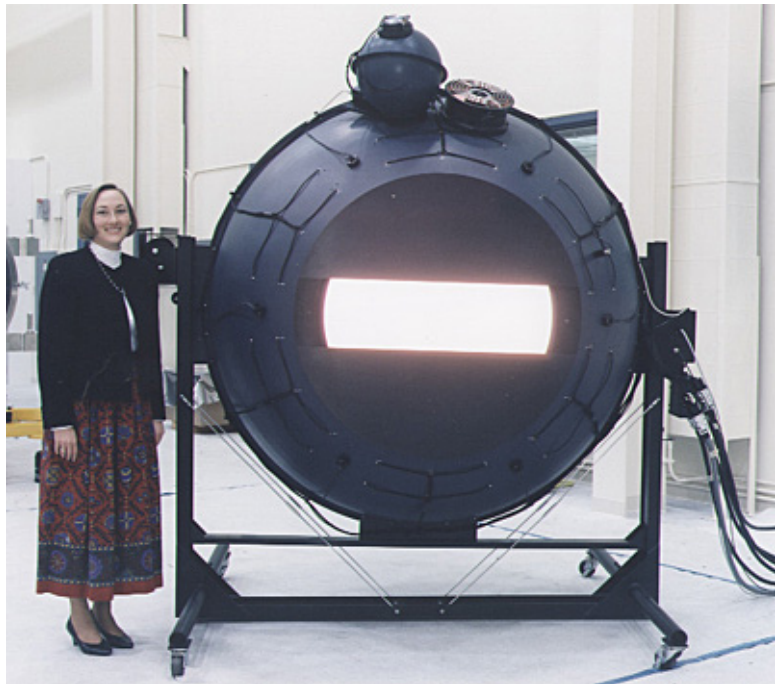
672 nm \pm 11 nm

866 nm \pm 20 nm

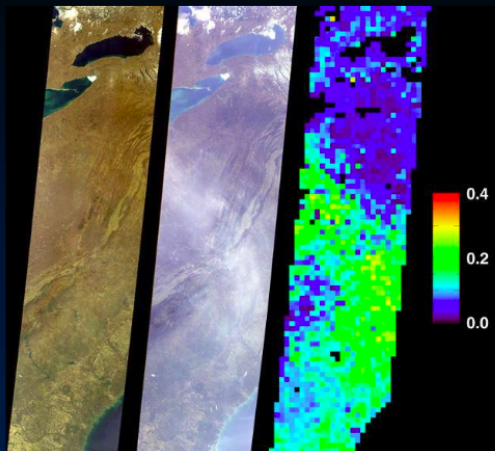


David J. Diner, JPL, Cal. Tech, Workshop May 22, 2005





1c. Enhancing sensitivity to thin aerosols



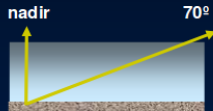
nadir

70°

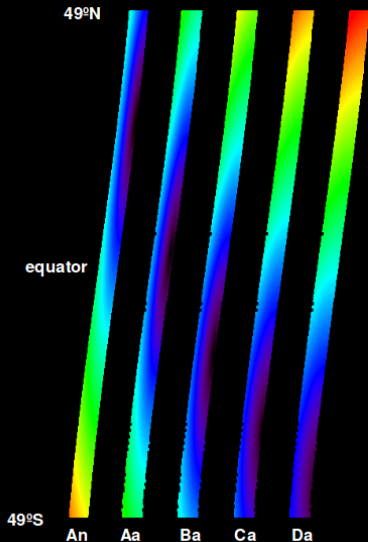
558-nm aerosol optical depth

Thin haze over land is difficult to detect in the nadir view due to the brightness of the land surface

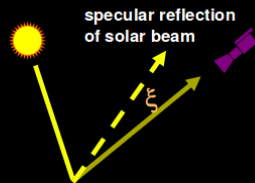
The longer atmospheric path length enhances the haze path radiance



Appalachians,
6 March 2000



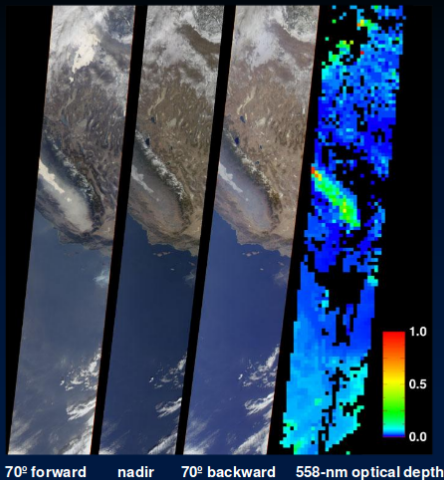
Example of glitter geometry July 3



Glitter
angle ξ

MISR aerosol
retrievals require
glitter avoidance
of at least 40°

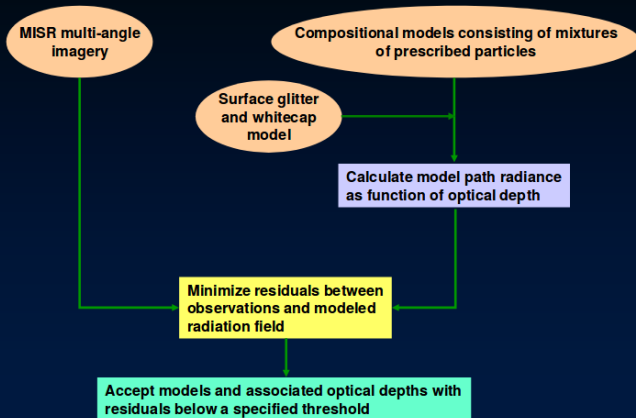
2a. Accounting for the surface contribution to TOA radiances



Despite the different ways of treating the surface, and the vast difference in water and land reflectance, good continuity is obtained across the land-water boundary

Southern California
and western Nevada
3 January 2001

Aerosol retrieval methodology over water



Multiple goodness of fit metrics

$$\chi^2_{\text{abs}} = N_{\text{channels}}^{-1} \sum_{\text{angle}} \sum_{\text{band}} [L_{\text{MISR}} - L_{\text{path}} - L_{\text{surface}}]^2 / [0.05 L_{\text{MISR}}]^2$$

where L_{surface} is modeled as a prescribed contribution from sunglint and whitecaps

χ^2_{geom}

: Similarly defined except measured and modeled radiances are normalized to the camera-average values

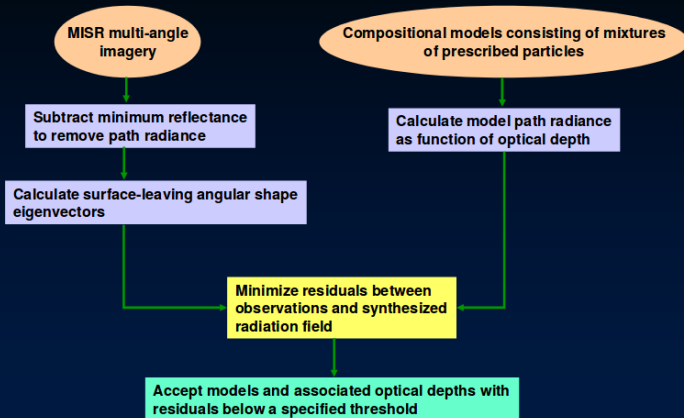
χ^2_{spec}

: Similarly defined except measured and modeled radiances are normalized to the red-band values

χ^2_{maxdev}

: Largest term in the χ^2_{abs} summation

Aerosol retrieval methodology over land



Goodness of fit metric

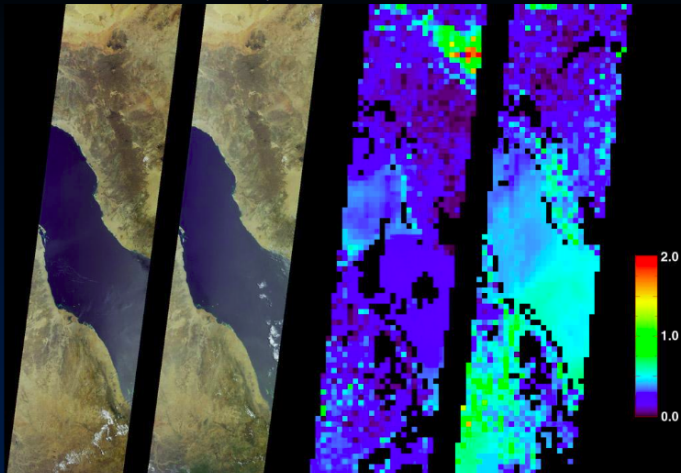
$$\chi^2_{\text{het}} = N_{\text{channels}}^{-1} \sum_{\text{angle}} \sum_{\text{band}} [L_{\text{MISR}} - L_{\text{path}} - L_{\text{surface}}]^2 / [0.05 L_{\text{MISR}}]^2$$

where L_{surface} is modeled as a dynamically derived sum of empirical orthogonal functions that are least-square fitted to $L_{\text{MISR}} - L_{\text{path}}$

Simplified concept:

- The technique requires surface contrast to be visible through the atmosphere
- Imagine two pixels with different albedos but the same variation in reflectance as a function of angle
- $L_{\text{MISR,TOA}(1)} = L_{\text{path}} + L_{\text{surface}(1)}$; $L_{\text{MISR,TOA}(2)} = L_{\text{path}} + L_{\text{surface}(2)}$
- $\Delta L_{\text{MISR,TOA}} = \Delta L_{\text{surface}}$ (path radiance subtracts out)
- The angular variation of L_{surface} is then given by $\Delta L_{\text{MISR,TOA}}$. To within a constant of proportionality, this is used to constrain $L_{\text{MISR}} - L_{\text{path}}$ by summing over all angles
- The EOF approach is invoked to account for multiple surface angular reflectance shapes within the scene

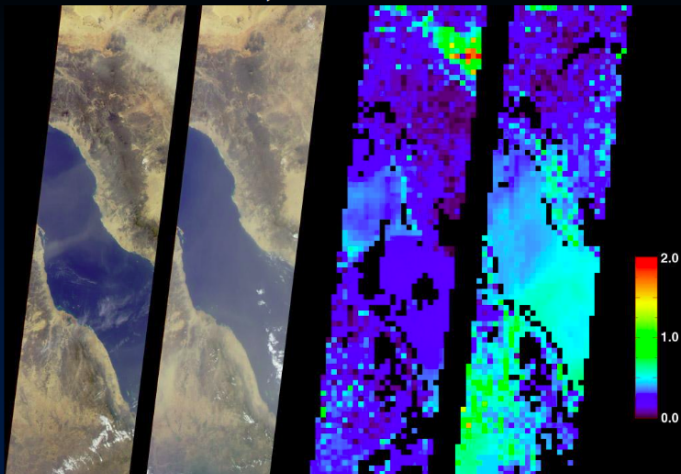
The Red Sea, 25 March and 29 June 2001



nadir images

558-nm aerosol optical depth

The Red Sea, 25 March and 29 June 2001



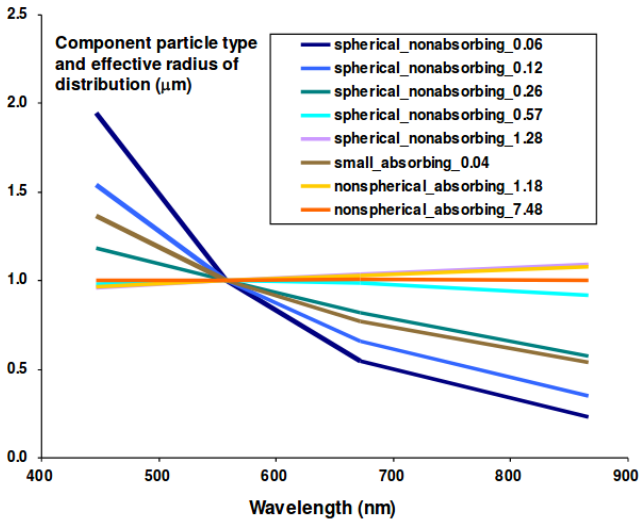
70°-forward images

558-nm aerosol optical depth

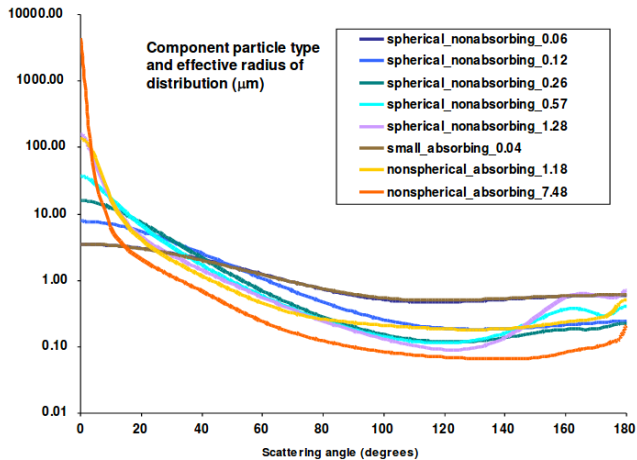
24 mixtures used in retrievals

		r_{eff} (components)	α	w_0
1	Spherical Small Clean	0.06	3.22	1.00
2	Spherical Small Clean	0.06, 0.12	2.71	1.00
3	Spherical Small Clean	0.12	2.24	1.00
4	Spherical Small Clean	0.12, 0.26	1.63	1.00
5	Spherical Medium Clean	0.26	1.09	1.00
6	Spherical Medium Clean	0.26, 0.57	0.56	1.00
7	Spherical Medium Clean	0.57	0.10	1.00
8	Spherical Medium Clean	0.57, 1.28	-0.05	1.00
9	Spherical Bimodal Clean	0.12, 1.28	0.82	1.00
10	Spherical Bimodal Clean	0.06, 1.28	1.19	1.00
11	Spherical Small Absorbing	0.06, 0.04	2.87	0.88
12	Spherical Small Absorbing	0.06, 0.12, 0.04	2.50	0.88
13	Spherical Small Absorbing	0.12, 0.04	2.09	0.88
14	Spherical Small Absorbing	0.12, 0.26, 0.04	1.62	0.88
15	Spherical Medium Absorbing	0.26, 0.04	1.13	0.88
16	Spherical Medium Absorbing	0.26, 0.57, 0.04	0.71	0.88
17	Spherical Medium Absorbing	0.57, 0.04	0.29	0.88
18	Dusty Low	0.26, 1.18	1.46	0.97
19	Dusty Low	0.26, 1.18	0.85	0.94
20	Dusty Low	0.26, 1.18	0.33	0.91
21	Dusty Low	1.18	-0.11	0.88
22	Dusty Low	1.18, 7.48	-0.08	0.83
23	Dusty Low	1.18, 7.48	-0.06	0.79
24	Dusty High	1.18	-0.11	0.88

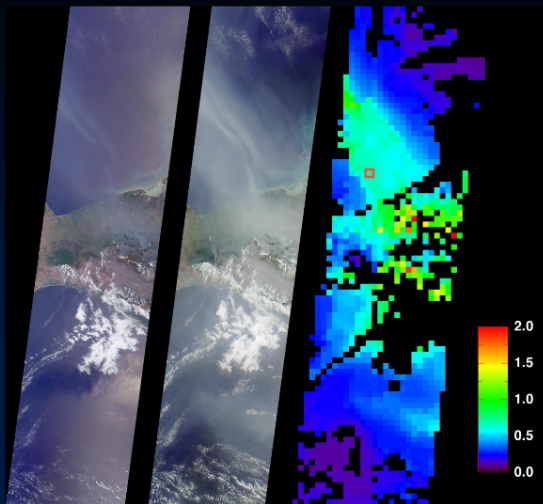
Spectral extinction of component aerosols relative to 558 nm



Scattering phase functions of component aerosols



Retrieval case study

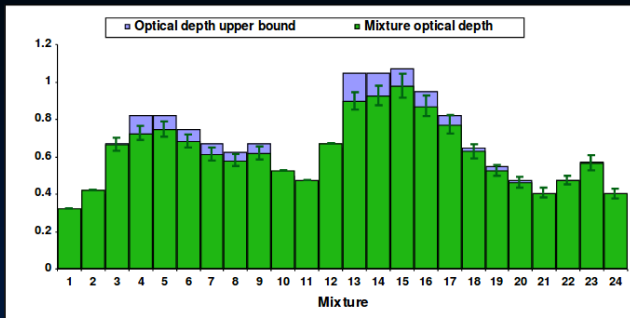


Southern Mexico
2 May 2002

558-nm optical depth

Retrieval case study

Orbit 12616, smoke

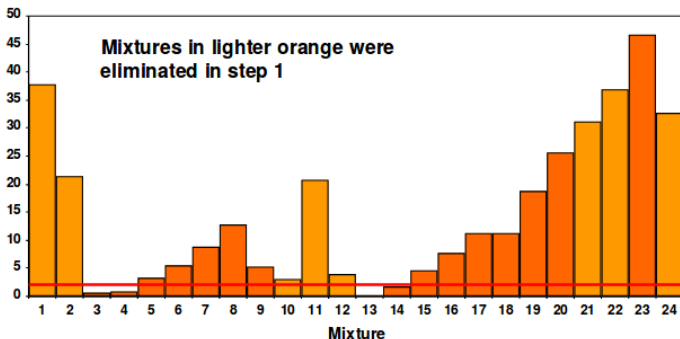


Optical depth is a function of aerosol type, so multi-angle and multi-spectral information is used to narrow the range of candidate solutions

Step 1: All 36 channels of MISR are used to establish an optical depth upper bound, and mixtures for which the best-fitting optical exceeds this limit are eliminated

Retrieval case study

Orbit 12616, smoke

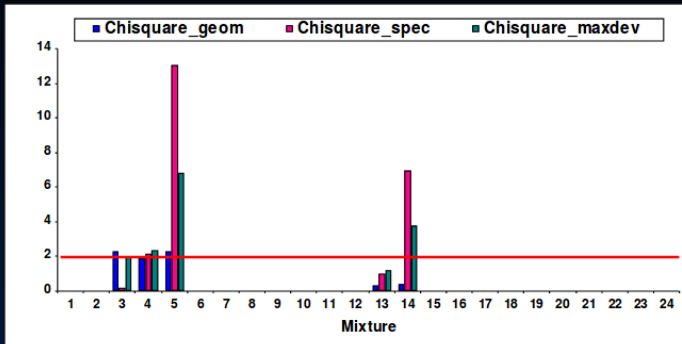


Step 2: Mixtures for which the χ^2_{abs} residual exceeds a specified threshold are eliminated

Ideally the threshold is ~1, but with quantized proportions of component particles in the mixtures, this is relaxed so as not to sacrifice coverage

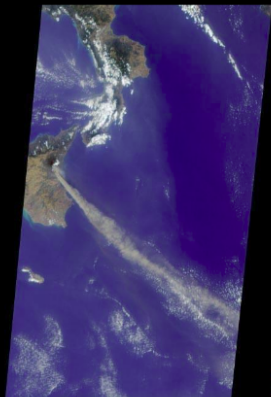
Retrieval case study

Orbit 12616, smoke

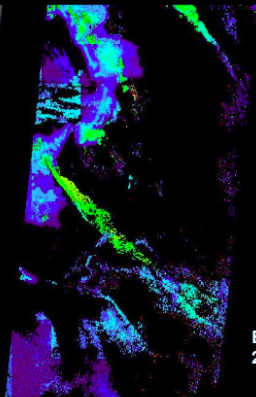


Step 3: Mixtures for which the other χ^2 residuals exceed specified thresholds are eliminated

For this case the best mixture is:
(13) Spherical Small Absorbing

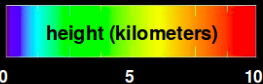


70° image

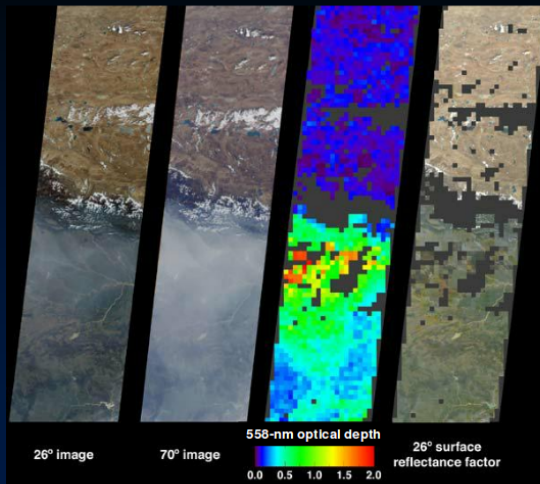


Automated
stereoscopic
retrieval
of plume height

Eruption of Mt. Etna, Sicily
22 July 2001



3c. Air quality



India and the
Tibetan Plateau
15 October 2001